

Gardiner Quadrangle, Maine

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Funding for the preparation of this map was provided in part by the U.S. Geological Survey STATEMAP Program, Cooperative Agreement No. 04HQAG0035.



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Open-File No. 09-8

2009

This map supersedes
Open-File Map 07-103.

SURFICIAL GEOLOGY OF MAINE

Continental glaciers like the ice sheet now covering Antarctica probably extended across Maine several times during the Pleistocene Epoch, between about 1.5 million and 10,000 years ago. The slow-moving ice superficially changed the landscape as it scraped over mountains and valleys, eroding and transporting boulders and other rock debris for miles. The sediments that cover much of Maine are largely the product of glaciation. Glacial ice deposited some of these materials, while others washed into the sea or accumulated in meltwater streams and lakes as the ice receded. Earlier stream patterns were disrupted, creating hundreds of ponds and lakes across the state. The map at left shows the pattern of glacial sediments in the Gardiner quadrangle.

The most recent "Ice Age" in Maine began about 25,000 years ago, when an icesheet spread southward over New England (Stone and Borns, 1986). During its peak, the ice was several thousand feet thick and covered the highest mountains in the state. The weight of this huge glacier actually caused the land surface to sink hundreds of feet. Rock debris frozen into the base of the glacier abraded the bedrock surface over which the ice flowed. The grooves and fine scratches (striations) resulting from this scraping process are often seen on freshly exposed bedrock, and they are important indicators of the direction of ice movement. Erosion and sediment deposition by the ice sheet combined to give a streamlined shape to many hills, with their long dimension parallel to the direction of ice flow. Some of these hills (drumlins) are composed of dense glacial sediment (till) plastered under great pressure beneath the ice.

A warming climate forced the ice sheet to start receding as early as 21,000 years ago, soon after it reached its southernmost position on Long Island (Sirkin, 1986). The edge of the glacier withdrew from the continental shelf east of Long Island and reached the present position of the Maine coast by 13,800 years ago (Dorion, 1993). Even though the weight of the ice was removed from the land surface, the Earth's crust did not immediately spring back to its normal level. As a result, the sea flooded much of southern Maine as the glacier retreated to the northwest. Ocean waters extended far up the Kennebec and Penobscot valleys, reaching present elevations of up to 420 feet in the central part of the state.

Great quantities of sediment washed out of the melting ice and into the sea, which was in contact with the receding glacier margin. Sand and gravel accumulated as deltas and submarine fans where streams discharged along the ice front, while the finer silt and clay dispersed across the ocean floor. The shells of clams, mussels, and other invertebrates are found in the glacial-marine clay that blankets lowland areas of southern Maine. Age dates on these fossils tell us that ocean waters covered parts of Maine until about 11,000 years ago,

when the land surface rebounded as the weight of the ice sheet was removed.

Meltwater streams deposited sand and gravel in tunnels within the ice. These deposits remained as ridges (eskers) when the surrounding ice disappeared. Maine's esker systems can be traced for up to 100 miles, and are among the longest in the country.

Other sand and gravel deposits formed as mounds (kames) and terraces adjacent to melting ice, or as outwash in valleys in front of the glacier. Many of these water-laid deposits are well layered, in contrast to the chaotic mixture of boulders and sediment of all sizes (till) that was released from dirty ice without subsequent reworking. Ridges consisting of till or washed sediments (moraines) were constructed along the ice margin in places where the glacier was still actively flowing and conveying rock debris to its terminus. Moraine ridges are abundant in the zone of former marine submergence, where they are useful indicators of the pattern of ice retreat.

The last remnants of glacial ice probably were gone from Maine by 10,000 years ago. Large sand dunes accumulated in late-glacial time as winds picked up outwash sand and blew it onto the east sides of river valleys, such as the Androscoggin and Saco valleys. The modern stream network became established soon after deglaciation, and organic deposits began to form in peat bogs, marshes, and swamps. Tundra vegetation bordering the ice sheet was replaced by changing forest communities as the climate warmed (Davis and Jacobson, 1985). Geologic processes are by no means dormant today, however, since rivers and wave action modify the land, and worldwide sea level is gradually rising against Maine's coast.

References Cited

- Davis, R. B., and Jacobson, G. L., Jr., 1985, Late-glacial and early Holocene landscapes in northern New England and adjacent areas of Canada: *Quaternary Research*, v. 23, p. 341-368.
- Dorion, C. C., 1993, A chronology of deglaciation and accompanying marine transgression in Maine: *Geological Society of America, Abstracts with Programs*, v. 25, no. 2, p. 12.
- Sirkin, L., 1986, Pleistocene stratigraphy of Long Island, New York, in Caldwell, D. W. (editor), *The Wisconsin stage of the first geological district, eastern New York*: New York State Museum, Bull. 455, p. 6-21.
- Stone, B. D., and Borns, H. W., Jr., 1986, Pleistocene glacial and interglacial stratigraphy of New England, Long Island, and adjacent Georges Bank and Gulf of Maine, in Sibbrava, V., Bowen, D. Q., and Richmond, G. M. (editors), *Quaternary glaciations in the northern hemisphere: Quaternary Science Reviews*, v. 5, p. 39-52.

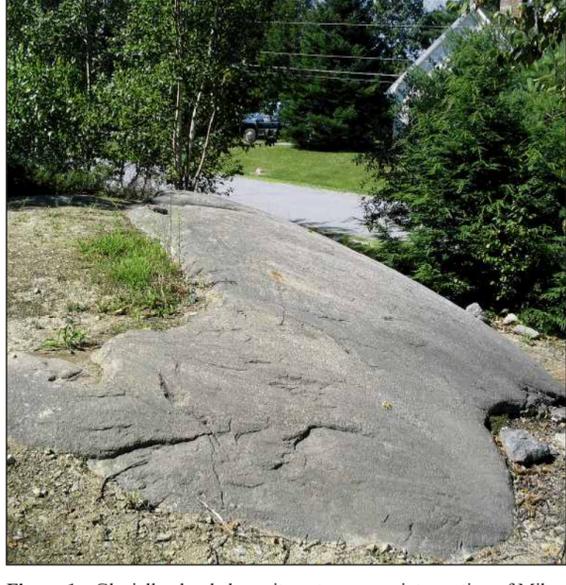


Figure 1: Glacially abraded granite outcrop near intersection of Mikes Lane and Maloy Avenue, north of Route 9/126 in West Gardiner. The asymmetric profile of the outcrop, with the gentler slope facing "up-glacier," and grooves on the ledge surface show that the glacier flowed east-southeast (right to left, as seen in photo).

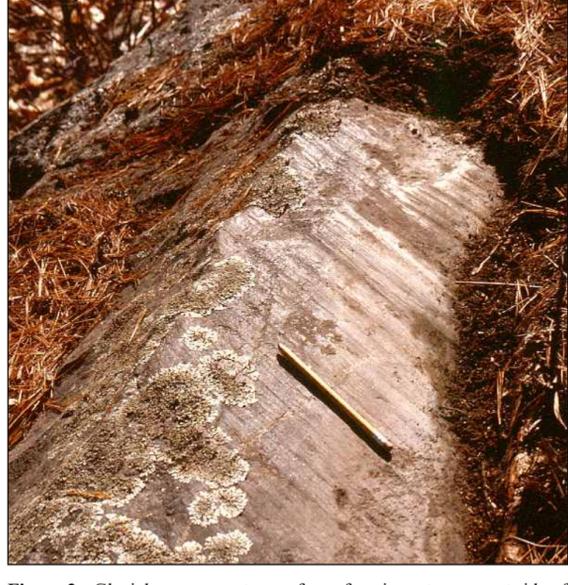


Figure 2: Glacial grooves on top surface of gneiss outcrop, west side of Kennebec River in South Gardiner. Pencil is parallel to ice flow direction (129°).



Figure 3: Pit exposure north of Beedle Road in Richmond, showing glacial till (above shovel) overlying and interlayered with deformed sand beds. The till was deposited at the base of the ice when it readvanced a short distance over the northern margin of a submarine fan.



Figure 4: View looking south-southeast across submarine fan in Richmond (same pit as Figure 3). The fan is composed of sand and gravel that washed into the sea at the edge of the last glacial ice sheet during its retreat from the area. A thin deposit of gray glacial-marine clay overlies the fan near left edge of photo.



Figure 5: Coarse gravel on west side of Kennebec River between Gardiner and South Gardiner. The esker was deposited by a meltwater stream in a subglacial ice tunnel. It is part of a long esker system that follows the Kennebec Valley, but much of this deposit is concealed under glacial-marine clay in the Gardiner area.



Figure 6: Gravel pit on east side of Kennebec River in Pittston, showing cross-section of esker ridge. The esker is composed of gravel, and it is directly overlain by gray glacial-marine clay. Unlike most pits in the Kennebec Valley, this exposure does not show a submarine fan deposit between the esker and younger clay unit.

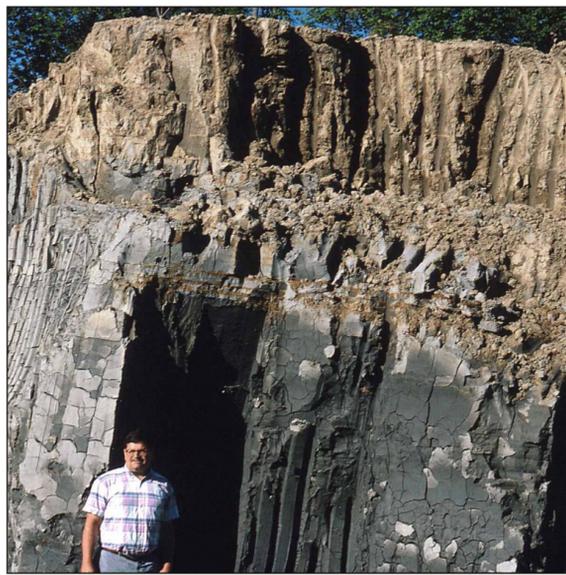


Figure 7: Glacial-marine clay (Presumpscot Formation) formerly exposed at construction site in Gardiner, near west end of Kennebec River bridge. Lower part of section shows fresh "blue clay," while the upper part has been oxidized to a brownish color.

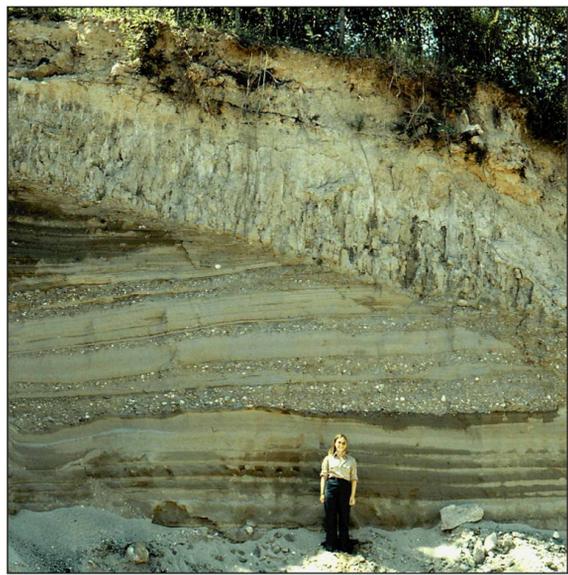


Figure 8: Pit face on east side of Kennebec River in Pittston, showing glacial-marine clay (Presumpscot Formation) overlain by sand and gravel (submarine fan). The latter unit was eroded by meltwater currents of a submarine landslide that truncated the eroded by beds prior to deposition of the clay.